

Stratigraphic Framework of the Wolfcamp – Spraberry of the Midland Basin

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Permian Basin Research Lab at UT Dallas

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Goals:

- Advance understanding of all geologic aspects of the Permian Basin through open applied research, linking academia and industry
- Educate and better prepare students for professional careers in the oil and gas industry
 - Graduate courses offered:
 - Geology of the Permian Basin
 - Petroleum Geoscience
 - Paleo Earth Systems: Global Themes

https://labs.utdallas.edu/permianbasinresearch/





- Review the tectono-stratigraphic framework of the Wolfcamp and Spraberry deep-water units of the Midland Basin, west Texas
- Briefly discuss the facies/characteristics of these rocks
- Highlight the differences between the Wolfcamp shale (A D) and Spraberry depositional systems

Note: although not specifically addressed, the framework outlined here is applicable to the Delaware Basin

Greater Permian Basin Region

 Confluence of Marathon-Ouachita fold and thrust belt and Ancestral Rockies basement-involved uplifts (Penn. – early Permian)



Tucumcari

SArbuck

OROGEN

Ardmore

Permian Basin Stratigraphy and Tectonic History



(modified from Reed, unpub., 2016)





A very dynamic time in Earth history, especially in west Texas

Wolfcampian – Early Leonardian

- Waning icehouse, transition to greenhouse
- Northward drift of Pangea
- Increasing aridity & expansion of continental desert in western U.S.
- Cratonic emergence / contraction of seaway (onset of long-term SL fall)
- Culmination of tectonic pulses in W. TX (mid WC); Pacific arc volcanism (Late WC-Leon.); PB enters rapid subsidence phase (Dean - Spraberry)

Late Pennsylvanian

- Icehouse climate; PB in humid-tropical setting (abundant rainfall)
- Numerous high-freq., high-amplitude sea-level changes
- Expansion of Penn seaway (long-term rise); stratified water columns
- Continued tectonism in west Texas (Marathon-Ouachita FTB, rise of ARM)





(maps: Ron Blakey, NAU/Colorado Plateau Geosystems)

Stratigraphic framework, Wolfcamp - Spraberry

Numerous 3rd- and higher-order cycles of sea-level change organized into larger 2nd-order trends (5 – 10+ m.y. in duration); from oldest to youngest:

WC D - lowermost WC C2
 WC C2
 WC C1
 WC A - B
 Spraberry - L. Clear Fork



FZ. Fusulinid zonation a. Tectonic pulses

b. Ash beds

c. Climate phase

mwu: mid-Wolfcamp unconformity

R = Regression T = Transgression

(Sea-level curve from Ross and Ross, 2009; Fusulinid zonation from Wahlman, 2019)

Midland Basin Type Log



Lower Strawn Limestone and Wolfcamp D (Cline)

STRAWN GROUP STRATIGRAPHY (DESMOINESIAN SERIES)

System Series Lithostratigraphic Unit Whitehorse Group; El Reno Group Guadalupian (part) Sumner Group, Enid Group, Hennessey Group Leonardian Permian Chase Group Council Grove Group Wolfcampian Pontotoc Group Admire Group Wabaunsee Group Ada Group Shawnee Group Virgilian Douglas Group Pennsylvanian Lansing Group Missourian Hoxbar Group Kansas City Group Marmaton Group Deese Group Desmoinesian Cherokee Grou Atoka Group Atokan Morrow Group/Formation Morrowan Springer Formation Mississippian Chesterian Chester Group Meramec lime Mayes Group Meramedian Osage lime Osadean Kinderhook Shale Kinderhookian Woodford Shale Chautauduan Devonian Misener sand Senecan Erian Ulsterian Cayugan Silurian Hunton Group Niagaran Alexandrian Sylvan Shale; Maguoketa Shale Cincinnatian Viola Group/Formation Ordovician Champlainian Simpson Group Canadian Arbuckle Group ambrian Trempealeauan Reagan Sandstone Franconian

Mid Continent / Anadarko Basin

(Gianoutsos et al, USGS, 2014)



Midland Basin Fusulinid zonation Late Late Strawn Middle Late Strawn } Marmaton

Early Late Strawn	
Late Middle Strawn Early Middle Strawn	Upper Cherokee
Late Early Strawn Middle Early Strawn Early Early Strawn	Lower Cherokee

Figure 1. Expanded fusulinid zonation of the Strawn, Hollingsworth approximate equivalents on the right. (Reed and Mazzullo, 1987)

Lower Strawn L	imestone = Early Strawn	

Middle & Upper Strawn = lower portion of WC D

AAPG COSUNA chart (1985)

Lower Desmoinesian Facies (Lower Strawn Limestone)



- Shallow water platform carbonate facies extend across entire Midland Basin and Eastern Shelf region
- Lower Strawn Limestone is generally
 200 ft. thick in Midland Basin
- Core analyses indicate typical Penn shelf cyclothem deposits: burrowed skeletal wackestones grading upward into phylloid algal packstones and skeletal grainstones, capped by exposure surfaces
- Pre-dates drowning of Midland, Delaware basins

(Wright, 2011)

Wolfcamp D (Canyon – Cisco) facies





(Saller et al., 1999)

Organic-rich Wolfcamp D (Canyon – Cisco) black shales in core from the center of Midland Basin



Wolfcamp D: Basinal cyclothems



- Equivalent to classic "Penn. cyclothems" on shelves
- Silica rich shales; relatively high clay content
- Each basinal cyclothem = 15 45 ft. thick; bounded by thin dolomite or LS; highly correlative basin-wide
- Organic content partitioned into multiple thin cycles
- High pore pressures due to depth, maturity



Wolfcamp C

Wolfcamp C

Lower Strawn

(Sinclair et al., 2018)

RDEEP N. Amer. GR Pioneer Preliminary correlation of Stage (API) ∆logR Zonation MB tops to Eastern Shelf 8200 2 West East Dean 8300 MIDLAND BASIN < EASTERN SHELF Eastern 8400 MB Shelf Mitchell Co. — Howard Co.-Taylor Co. - Nolan Co.-8500 A2 Wolfcamp A ∢ 8600 Elm Creek B1 Admiral-മ Coleman Junc 8800 B2 Group 8900 Sedwick--WFMP A-B Ó Wolfcamp B Ibex Noodle 9100 Noodle Creek Cisco PBY - Dean Camp Creek 9200 20 WEMPC Saddle Creek-Crystal Falls MP C2 Mid Unc (modified from WTGS, 1984, Breckenridge-Wolfcamp C1 Finis Sh. 9500 and Brown et al., 1990) WFMP 9600 Canyon Gp. 9700 M. – U. Strawn 9800 9900 Westward progradation of Eastern Shelf delta Lower Strawn Wolfcamp C2 ٠ 1000 systems and platform margins (100 -150 km) Miss-ourian Initial development of Glasscock Nose during WFMP C1 time Wolfcamp D

• Uplift of CBP structural blocks and development of mid-Wolfcamp unconformity

Progradation of Wolfcamp C shelf delta systems across the Eastern Shelf



Sequential development of the Glasscock Nose





	Period	Stage	N. Amer. Stage	Glass Mountains	Central Basin Platform		Midland Basin		
I	ſ	Kungur	_	Cathedral	Holt / Upper Leonard	Clear- fork	Holt / Upper L	eonard	
			ian	ian	Mountain	Glorieta	È	Upper	
			Skinner Ranch / Hess	Upper Clearfork / Yeso	ber	Middle			
		Artin- skian Sakmar- ian Artin- skian Sakmar- ian Sakmar- ian Sakmar- ian Skinner Ranch / Hess Lenox Hills (Upper Wolfcamp				Jo M			
	iaı			Middle Clearfork / Yeso	0)	L. Spra	b. Sh		
	εl			Lwr Clearfork / Yeso		Wolfcamp	A		
nnsyl- Per	Per		Wolfcamp	Wolfcamp B					
			(Upper Wolfcamp)	Konto and a second seco	Wolfcamp C1				
		Assel- ian	Assel- ian Year Ranch (Wr Wolfcamp" ur	ndonfarmity	1	Wolfcamp	C2		
	ısyl- ian	Gzhelian Kasi- Missourian Moscovian Besmoinesian	Cisco		Cisco	Wolf-			
	an		Canyon	Canyon					
	Pe v		Strawn		Strawr	<u></u>			

- last major tectonic pulse prior to middle – late Permian subsidence phase
- note diachronous nature of unconformity across Permian Basin region

Wolfcamp A - B

Wolfcamp A and B Facies & Depositional Model



Wolfcamp A - B



- 700+ ft. of organic-rich, silica- and calcareous-rich mudstone punctuated by numerous density flows (carb. turbidites and debris flows)
- Six operational sub-units:
 - A1 B1
 - A2 B2
 - A3 B3
- WC B are predominantly siliceous mudstones
- WC A are mixed carb-silica mudstones
- Aggradation of carbonate margins during second-order highstand increase percentage of CaCO₃ into basin during WFMP A time
- Interval currently resides in peak oil window in Midland Basin; remains a main horizontal drilling target





Wolfcamp B2



Wolfcamp A3



Photograph of core, **Wolfcamp B2**, depth 8837-8847 feet. (A) Structureless silty mudstone with phosphatic concretion. (B) Calcareous silty mudstone. (C) Carbonate lithoclast. (D) Ash bed. (E) Carbonate concretion. (F) Skeletal grainstone with erosive base and reworked concretions. (G) Thin, muddy debrite with deformed mudclast. (H) Sheared and rotated package of thin beds at the bottom of a slumped interval, 8847-8843 ft.

(Murphy, 2105)

Wolfcamp carbonate debris flows





- Flows are thickest and coarsest near the shelf margins; distal portions of flows are thinner and finer grained
- Geometries include sheet-like fans and highly channelized flows

Spraberry - Dean

Midland Basin Type Log









Porous

sandstone

Sandstone

ferroan

dolomite

cemented w/

(a)

(b)

Laminated siltstones

- All fans (Dean, Jo Mill, Middle & Upper Spraberry) are similar in appearance
- Main facies:
 - Massive f.g. sandstones ("Bouma A")
 - Laminated siltstones / shales
 - Burrowed siltstones / shales (O₂)
 - Black shale (thin caps)
- Provenance? (north vs. south)
 Depositional model ?



Strong age signal ("peaks")

- Grenville province (1100 Ma)
- Gondwana (600 Ma)
- Appalachia (400 Ma)
- Intermediate signal
 - Granite-Rhyolite province (1400 Ma)
- Weak signal
 - Yavapai Mazatzal province (1700 Ma)

Strong age signals are from southern-located provinces, indicating a southern source land for Dean – Spraberry sands in central & southern Midland Basin (currently accepted view: all sands were from a northern source)

(Waite et al., in press)

Spraberry & Dean (Bone Spring) Depositional Model (based on Hanford, 1981)

Spraberry and shelf equivalents are alternating sand-rich and organic shale/carbonate-rich packages deposited during alternating high and low sea levels.



Organic-rich shales highly laminated and not bioturbated; Organic-poor shales bioturbated Thin dolomitic hard grounds observed in sands and shales





Highstand -

- Shelf submerged
- · Carbonates on shelf
- Carbonate gravity flow deposits and organic-rich shales in basin

Lowstand and ensuing transgression-

- Shelf exposed
- · Clastics move across shelf via wind and in wadis
- Clastic gravity flow deposits bypass shelf during lowstand and are cannabalized during early transgression

Possible modern analog for Dean - Spraberry: Offshore Mauritania, African Sahara





Summary and Conclusions

- The Wolfcamp Spraberry interval of the Midland Basin consists of a series of lithologically- and mineralogically-complex facies; each interval is unique
 - Wolfcamp D: basinal cyclothems
 - Wolfcamp C: clay-rich shales
 - Wolfcamp A B: Silty, calcareous terrigenous shales; carbonate % increases upward
 - Dean Spraberry: Argillaceous siltstones, punctuated by numerous submarine-fan complexes (massive & laminated sandstones)
- Complexity of these rocks reflects changing/evolving geologic conditions (eustasy, climate, tectonics, sediment supply, biota, etc.) along the SW margin of western Pangea during Late Pennsylvanian – early Permian time
- Geologists must work closely with drilling, completion, and reservoir engineers to fully communicate the complexity and uniqueness of each unit / horizontal zone

"Not all shales are created equal"



(Hamlin and Baumgardner, 2012)